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Metastable equilibria among dicarboxylic acids and the oxidation state during aqueous alteration on the CM2 chondrite parent body

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Abstract

Linear saturated dicarboxylic acids are present in carbonaceous chondrite samples at concentrations that suggest aqueous alteration under conditions of metastable equilibrium. In this study, previously published values of dicarboxylic acid concentrations measured in Murchison, Yamato-791198, and Tagish Lake carbonaceous chondrites are converted to aqueous activities during aqueous alteration assuming water:rock ratios that range from 1:10 to 10:1. Logarithmic plots of the aqueous activities of any two dicarboxylic acids are proximal to lines whose slope is fixed by the stoichiometry of reactions describing the oxidation–reduction equilibrium between the two species. The precise position of any line is controlled by the equilibrium constant of the reaction relating the species and the hydrogen fugacity for the reaction of interest. Reactions among succinic (C4), glutaric (C5), and adipic (C6) acids obtained from CM2 chondrites show evidence of metastable equilibrium and yield $\log f_{\text{H}_2}$ values that agree to within 0.3 log units at 298.15 K and 0.6 log units at 473.15 K. At a water:rock ratio of 1:1, metastable equilibrium among succinic, glutaric, and adipic acids results in calculated $\log f_{\text{H}_2}$ values during aqueous alteration that range from -6.2 at 298.15 K to -3.3 at 373.15 K. These values are consistent with those obtained in previous work on carbonaceous chondrites and with metastable equilibrium at temperatures ranging from 300 to 355 K in contact with cronstedtite + magnetite.

1. Introduction

Carbonaceous chondrites represent a subset of meteorites from which an abundant suite of organic compounds have been extracted (reviewed by Botta and Bada, 2002; Sephton, 2002; Pizzarello et al., 2006). There is considerable interest in the origin of the reduced carbon species present in these meteorites: these species represent a potential initial stock of organic compounds for the early Earth (Chyba and Sagan, 1992). The carbonaceous chondrites are derived from multiple asteroidal parent bodies that underwent post-accretionary aqueous alteration (Barber, 1985; Kerridge, 1991; Krot et al., 2003). Although the timing and mechanism of organic synthesis remains unknown, one possibility is that organic synthesis was coincident with

all or part of the aqueous alteration (Kerridge and Bunch, 1979; Barber, 1985; Cronin et al., 1993). Schulte and Shock (2004) have tested that hypothesis by simulating organic synthesis under conditions of metastable equilibrium during meteorite alteration. Their work indicates that synthesis of classes of hydrophilic organic species during aqueous alteration was possible. It is appropriate therefore to consider whether evidence of alteration of organic species or organic synthesis is preserved in the abundances of individual species within these compound classes.

The dicarboxylic acids are a particularly apposite choice for this examination. They are simple chemically, consisting only of carbon, hydrogen, and oxygen arranged in simple functional groups, whereas other organic compounds found in meteorites have more complex elemental or struc-

tural chemistry. The presence of a carboxylic acid functional group on each end of the molecule suggests that the terminal carbon atom experienced electron transfer: a free-radical termination reaction involving two carboxylic acid radical species in a dilute aqueous solution is unlikely.

The synthesis of organic compounds found in carbonaceous chondrites is an area of debate and continued research; "given the large compositional and isotopic differences (among the organic compounds identified in carbonaceous chondrites) an inclusive theory for the origin of all the varied organic constituents in carbonaceous chondrites appears hard to formulate" (Pizzarello et al., 2006). Organic compounds isolated from carbonaceous chondrites exhibit large D/H ratios and positive $\delta^{13}\text{C}$ values, indicating that the first organic molecules, or at least the reactants leading to their production via proposed Fisher-Tropsch type (FTT) (Anders et al., 1973) or Strecker (Peltzer et al., 1984) synthesis pathways, were the products of low temperature ion reactions in an interstellar environment (Cronin et al., 1993; Pizzarello and Huang, 2002). These molecules were analogous to the anhydrous inorganic mineral species that existed on the parent body before alteration. Irrespective of the mechanism of synthesis, organic compounds were present on the parent body and hydrophilic species occurred as solutes in the aqueous solution during the alteration event (Brearley, 2006; Krot et al., 2006). Additional hydrophilic organic molecules may have been synthesized from oxidized forms of carbon during aqueous alteration whereas others were formed from preexisting hydrophobic or hydrophilic organic species. The suite of resulting compounds is a product of pre-accretionary synthesis and the aqueous alteration event. The chemistry of the latter event, and perhaps the distribution of the hydrophilic organic species, was controlled partly by the f_{H_2} environment imparted by the mineral assemblage.

This work attempts to test the evidence for metastable equilibrium among dicarboxylic acids extracted from carbonaceous chondrites. Stable thermodynamic equilibrium predicts that carbon containing compounds will exist either in the oxidized form of CO_2 or the reduced forms of CH_4 , graphite, or diamond. The abiotic synthesis and preservation of organic compounds in carbonaceous chondrites would not be expected based on stable thermodynamic equilibrium. Minimizing chemical potential requires favorable reaction kinetics: the reaction products favored thermodynamically will only be obtained provided there is a minimal activation energy to allow the reaction to occur. It is, however, possible for some systems to achieve metastable equilibrium. Reactions at metastable equilibrium are reversible and the relative abundance of species in the reaction of interest are fixed by the equilibrium constant of the reaction. Metastable equilibrium has been documented for reactions between ethane and ethene (Seewald, 1994) and for the interconversion of formic acid to carbon dioxide (McCullom and Seewald, 2003). There is evidence that acetic acid and propanoic acid, and propanoic acid and butanoic acid are in metastable equilibrium in sedimentary basins (Shock, 1988, 1989, 1990; Helgeson et al., 1993). Thus the possibility that speciation of dicarboxylic acids during aqueous alteration on asteroids was controlled by conditions of metastable equilibrium

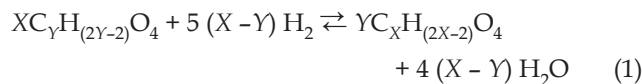
is an intriguing one. If evidence of metastable equilibrium among the dicarboxylic acid species can be obtained, then the oxidation state of the aqueous alteration event during which the dicarboxylic acids equilibrated can be retrieved. The calculated oxidation state would allow estimates of the oxidation state of the fluid during aqueous alteration of the parent body to be constrained.

2. Methods

Five studies of dicarboxylic acids in carbonaceous chondrites were reviewed from literature (Table 1). Dicarboxylic acid concentrations have been measured in (1) Murchison (CM2) and (2) Yamato-791198 (CM2) meteorites (Shimoyama and Shigematsu, 1994) following extraction with water at 298 K (A. Shimoyama, private communication, 2006). High-temperature extraction of dicarboxylic acids was performed for the (3) Tagish Lake (Unclassified, exhibiting CM to CI characteristics; Zolensky et al., 2002) meteorite (extraction in water for 24 h at 383 K) (Pizzarello and Huang, 2002), (4) Murchison meteorite (extraction in water for 20 h at 373 K) (Peltzer, 1979), and (5) Murchison meteorite (extraction in water for 24 h at 373 K) (Martins et al., 2006).

Assessment of metastable equilibria based on the relative concentrations of organic acids follows the methods of Shock (1988, 1989). The slopes of log activity plots of organic acids involved in redox reactions are shown to correspond to those predicted by the linear rearrangement of the mass action equation governing the reaction of interest. Use of this method requires the calculation of the activities of dicarboxylic acids present during aqueous alteration on the asteroid. Measured concentrations (nmol/g meteorite) (Table 1) were converted to molal concentrations (mol/kg solvent) at water:rock mass ratios of 1:10, 1:1 and 10:1 and then to aqueous activities (Table 1) by assuming unit activity coefficients (γ). These water:rock ratios bracket those proposed for CM chondrite alteration (Clayton and Mayeda, 1999) and are consistent with models of organic synthesis during aqueous alteration on asteroids that have utilized water:rock ratios of 1:1 to 10:1 by mass (Schulte and Shock, 2004).

Consideration of solubility data (Seidell, 1941; Apelblat and Manzurola, 1987) suggests that the concentrations of dicarboxylic acids measured in carbonaceous chondrites are sufficiently low that these species would have been undersaturated during aqueous alteration on the asteroid and during extraction. Reactions among linear saturated dicarboxylic acids have the form:



where X and Y are stoichiometric reaction coefficients fixed by the selection of the species of interest. Linear rearrangement of the mass action equation yields:

$$\log a_{\text{C}_Y\text{H}_{(2Y-2)}\text{O}_4} = (Y/X) \log a_{\text{C}_X\text{H}_{(2X-2)}\text{O}_4} - (1/X) [5(X - Y) \log f_{\text{H}_2} + \log K_r] \quad (2)$$

Table 1. Measured concentrations (nmol/g meteorite) of dicarboxylic acids and corresponding activities (*a*) calculated based on a water:rock (W:R) ratio of 1:1

Meteorite	C2 (Oxalic) nmol/g	Log <i>a</i>	C3 (Malonic) nmol/g	Log <i>a</i>	C4 (Succinic) nmol/g	Log <i>a</i>	C5 (Glutaric) nmol/g	Log <i>a</i>	C6 (Adipic) nmol/g	Log <i>a</i>	C7 (Pimelic) nmol/g	Log <i>a</i>	C8 (Suberic) nmol/g	Log <i>a</i>	C9 (Azelaic) nmol/g	Log <i>a</i>
Murchison ^a	2100	-2.7	96	-4.0	97	-4.0	24	-4.6	5.8	-5.2	2	-5.7	1.7	-5.8	3.2	-5.5
Yanato-791198 ^a	1900	-2.7	20	-4.7	21	-4.7	6.8	-5.2	2.3	-5.6	0.5	-6.3	0.3	-6.5	0.3	-6.5
Tagish Lake ^b	N/D	N/D	3	-5.5	40	-4.4	28	-4.6	19	-4.7	1.5	-5.8	<1	<-6	<1	<-6
Murchison ^c	102	-4.0	12.5	-4.9	50	-4.3	25	-4.6	4.8	-5.3	N/D	N/D	N/D	N/D	N/D	N/D
Murchison ^d	N/D	N/D	N/D	N/D	54.2	-4.3	1.3	-5.9	1.8	-5.7	0.4	-6.4	0.2	-6.8	0.1	-6.8

N/D, not determined.

$\text{Log } a_{(\text{W:R}=10:1)} = \text{log } a_{(\text{W:R}=1:1)} - 1$.

$\text{Log } a_{(\text{W:R}=1:10)} = \text{log } a_{(\text{W:R}=1:1)} + 1$.

^a nmol/g data: Shimoyama and Shigematsu (1994).

^b nmol/g data: Pizzarello and Huang (2002).

^c nmol/g data: Peltzer et al. (1984).

^d nmol/g data: Martins et al. (2006).

where *a* represents the activity of the subscripted species, *f* is the fugacity, and K_r is the equilibrium constant for the reaction in Equation (1). This rearrangement yields the equation for a straight line in $\text{log } a_{\text{C}_{\text{YH}}(\text{2Y-2})\text{O}_4}$ vs. $\text{log } a_{\text{C}_{\text{XH}}(\text{2X-2})\text{O}_4}$ space. If two dicarboxylic acids are in metastable equilibrium, then a plot of the logarithm of the activities of these two species will yield a straight line with a slope (*Y/X*) fixed by the stoichiometry of the reaction and an intercept (*b*)

$$b = - (1/X) [(5(X - Y) \text{log } f_{\text{H}_2} + \text{log } K_r) \quad (3)$$

fixed by the equilibrium constant (K_r), the reaction stoichiometry, and the oxidation state ($\text{log } f_{\text{H}_2}$).

The value of $\text{log } f_{\text{H}_2}$ is calculated from the intercept by rearranging Equation (3):

$$\text{log } f_{\text{H}_2} = (-bX - \text{log } K_r) / [5(X - Y)]. \quad (4)$$

Because the slope of the line that corresponds to Equation (2) is fixed by the reaction stoichiometry it is inappropriate to generate the slope and intercept by a conventional least squares analysis. The $\text{log } a_{\text{C}_{\text{YH}}(\text{2Y-2})\text{O}_4}$ -intercept was therefore calculated by minimizing the variance (s^2) which is calculated according to Equation (5)

$$s^2 = \sum (y_i - y_{\text{calculated}})^2 \quad (5)$$

y_i is the observed value of the $\text{log } a_{\text{C}_{\text{YH}}(\text{2Y-2})\text{O}_4}$ and $y_{\text{calculated}}$ is the value of $\text{log } a_{\text{C}_{\text{YH}}(\text{2Y-2})\text{O}_4}$ given the requisite slope and the calculated intercept. The intercept is then calculated from

$$b = [\sum y_i - (Y/X) \cdot \sum x_i] / N \quad (6)$$

where x_i is the value of $\text{log } a_{\text{C}_{\text{XH}}(\text{2X-2})\text{O}_4}$ and *N* is the number of observations. Because the value of the variance is dependent on *N*, division by *N* allows the variance to be normalized and data sets of different sample sizes to be compared.

Note that in minimizing the variance, the slope of the lines does not indicate a best fit to the data; rather the resulting line indicates log activity values predicted by metastable equilibrium between two compounds. The proximity of the data to this line is indicated by the variance, allowing for quantitative comparison among data sets. While the selection of the water:rock ratio will not affect the slope of the activity plots, it will affect the value of the $\text{log } a_{\text{C}_{\text{YH}}(\text{2Y-2})\text{O}_4}$ -intercept. To assess the sensitivity of this factor, calculations were made with water:rock ratios of 10:1, 1:1, and 1:10 by mass.

Calculation of the $\text{log } a_{\text{C}_{\text{YH}}(\text{2Y-2})\text{O}_4}$ -intercept allows the fugacity (*f*) of H_2 at a given temperature (Equation (3)) to be retrieved provided the equilibrium constant is known. Equilibrium constants for the reactions of interest (K_r) (Table 2) were obtained from the change in standard Gibbs energy for the reaction

$$\text{log } K_r = -\Delta_r G^\circ / (2.303RT). \quad (7)$$

Values of the standard Gibbs energy of formation ($\Delta_r G^\circ$) for each acid of interest were calculated via a functional group contribution model (Yezdimer et al., 2000) at the saturation pressure of water. The functional group additivity model of Yezdimer et al. (2000) assigns an equal contribution to a "hydrocarbon" carbon atoms irrespective of whether the carbon atom is a primary, secondary, or tertiary carbon.

Table 2. Equilibrium constants for reactions between succinic, glutaric, and adipic acids at 295.15–473.15 K at the saturation pressure of water

T (K)	log K_r					
	A	B	C	D	E	F
298.15	33.0	66.0	-33.0	33.1	-66.0	-33.1
348.15	22.8	45.3	-22.8	22.4	-45.3	-22.4
373.15	18.8	37.2	-18.8	18.3	-37.2	-18.3
423.15	12.5	24.3	-12.5	11.7	-24.3	-11.7
473.15	7.5	14.3	-7.5	6.5	-14.3	-6.5

Reactions:

- (A) $5C_4H_6O_4 + 5H_2 = 4C_5H_8O_4 + 4H_2O$;
 (B) $6C_4H_6O_4 + 10H_2 = 4C_6H_{10}O_4 + 8H_2O$;
 (C) $4C_5H_8O_4 + 4H_2O = 5C_4H_6O_4 + 5H_2$;
 (D) $6C_5H_8O_4 + 5H_2 = 5C_6H_{10}O_4 + 4H_2O$;
 (E) $4C_6H_{10}O_4 + 8H_2O = 6C_4H_6O_4 + 10H_2$;
 (F) $5C_6H_{10}O_4 + 4H_2O = 6C_5H_8O_4 + 5H_2$.

Although this approach yields a model that can be applied to a wide variety of organic compounds, it also results in the calculation of identical values for thermodynamic properties of isomeric compounds. Calculation of $\Delta_f G^\circ$ for isomeric compounds using the Yezdimer et al. (2000) model is, therefore, dependent on accurate experimental values of the enthalpy ($\Delta_{hyd} H^\circ$) and entropy of hydration ($\Delta_{hyd} S^\circ$) of the compound of interest at 298.15 K. These values are not available, however, for many branched dicarboxylic isomers of linear dicarboxylic acids.

Thermodynamic calculations for inorganic species were made using The Geochemist's Workbench v. 5.0.2 and the thermo.com.v8.r6+.dat database.

3. Results

The plots in Figure 1 are used to evaluate the evidence that the C4–C9 linear saturated dicarboxylic acids were in metastable equilibrium during aqueous alteration on the parent body. The fill patterns of the symbols in Figure 1 represent the individual dicarboxylic acids in this study. The shapes of the symbols in Figure 1 represent the carbonaceous chondrite from which the acid was extracted. To prevent congestion on the plots, an offset (Z) is incorporated wherein the values for the log activities of suberic (C8) and azelaic (C9) acids are decreased by 0.5 and 1.0 log units, respectively. No offset is required for the other acids and thus $Z = 0$ for these species. Each line is labeled with the normalized variance (s^2/N) associated with the data points obtained from the CM2 chondrites (Figure 1).

All known measurements of dicarboxylic acids were included in Table 1. The results from the Tagish Lake samples are generally similar to those obtained from the CM2 chondrite samples (Table 1 and Figure 1). It has been suggested that CM and CI carbonaceous chondrites originated from progressive alteration of similar parent body assemblages (Zolensky et al., 1989) and that stones classified as CM and CI, as well as Tagish Lake, may have originated from the same parent body (Brandon et al., 2005). Whereas the $\Delta^{17}O$ values of carbonates and structural and inter-layer water from the Murchison meteorite are negative, the corresponding values from the Tagish Lake meteorite

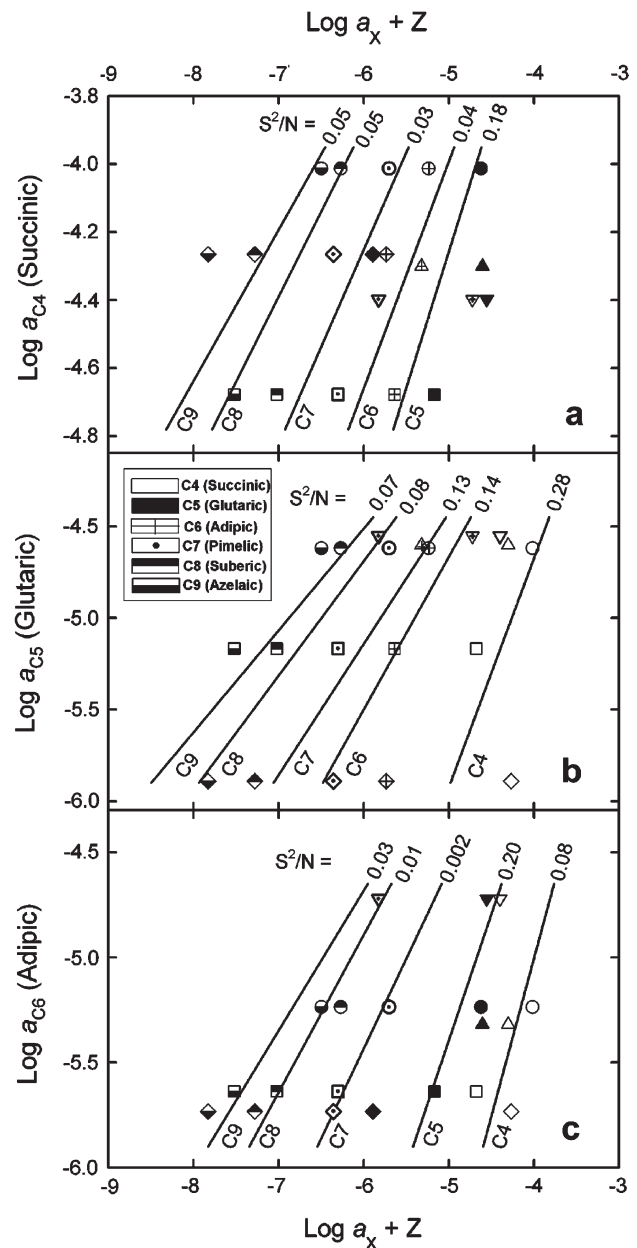


Figure 1. Plots of (a) $\log a_{C4}$ (Succinic), (b) $\log a_{C5}$ (Adipic), (c) $\log a_{C6}$ (Glutaric) versus $\log a$ linear saturated dicarboxylic acids of 4–9 carbon atoms reveal trends along lines of slope indicating metastable equilibrium among the acids. The normalized variance is indicated in bold above each line. The shape of the plot indicates data source: (○) Murchison (Shimoyama and Shigematsu, 1994), (□) Yamato-791198 (Shimoyama and Shigematsu, 1994), (▽) Tagish Lake (Pizzarello and Huang, 2002), (Δ) Murchison (Peltzer et al., 1984), (◇) Murchison (Martins et al., 2006). For suberic (C8) and azelaic (C9) acids Z equals -0.50 and -1.0 , respectively; Z equals zero for all other acids. The plots represent activities calculated based on a water:rock ratio of 1:1. The slope of the lines and the calculated variance is independent of the water:rock ratio.

are positive (Baker et al., 2002). Thus alteration of the CM2 chondrites and the Tagish Lake meteorite occurred in sufficiently different environments that dicarboxylic acids from those meteorites could not have been in metastable equilibrium even if the objects were derived from the same parent body. Moreover, removal of the Tagish Lake samples does

improve the fit of the data. If the Tagish Lake samples are removed the values of s^2/N decrease for seven of the nine possible permutations (Figure 2). The Tagish Lake samples have been plotted in Figure 1 for reference only and have not been used in calculating the positions of the lines or the values of the normalized variance (s^2/N) that are presented in Figure 1.

The slopes of the lines in Figure 1(a) are 4/5, 4/6, 4/7, 4/8, and 4/9 for the lines labeled C5, C6, C7, C8, and C9, respectively. Although the slopes are fixed, the position of each line is variable and is set by the intercept. The slopes predicted by Equation (2) are, however, relevant to the fitting of the plots. For example, in Figure 1c, the plots of $\log a_{\text{Adipic (C6)}}$ versus $\log a_{\text{Azelaic (C9)}}$ fit well ($s^2/N = 0.03$) to the slope of 6/9 as required by the reaction stoichiometry. Use of another value for the slope would provide a poorer fit. For example, use of a slope of 6/4 for these same data would yield a normalized variance with a much higher value ($s^2/N = 0.41$). The values of normalized variance obtained for the CM2 chondrite samples (Figure 1) compare favorably to normalized variance values of ~ 0.02 to 0.05 (calculated using data from Carothers and Kharaka, 1978; Willey et al., 1975) for the plots used by Shock (Shock, 1988, Figure 5; Shock, 1990, Figures 2–4) to demonstrate metastable equilibrium within oil field waters and sedimentary basins.

The reported concentrations of oxalic (C2) and malonic (C3) acids range widely among the four samples (Table 1). Malonic acid decarboxylates at low temperatures relative to other dicarboxylic acids; with a half life of approximately 3.5 h at 373 K (Maiella and Brill, 1996). Significant amounts of malonic acid would have decarboxylated during extraction of Murchison at 373 K for 20 h (Peltzer, 1979) and extraction of Tagish Lake at 383 K for 24 h (Pizzarello and Huang, 2002). The reason for the inconsistencies in reported oxalic acid concentrations is not readily apparent; non-catalytic decarboxylation of oxalic acid is minor over the extraction time intervals (Weber, 2004).

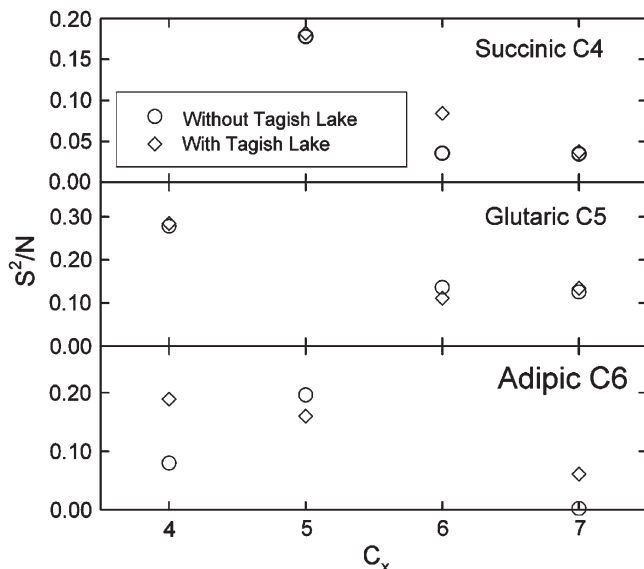


Figure 2. Plots of the normalized variance for fits both including and excluding results from the Tagish Lake meteorite.

Variations in dicarboxylic acid concentrations among Murchison samples may be a function of sample heterogeneity, extraction, and/or analysis techniques. Note that relevance lies in the ratios among dicarboxylic acids extracted from a sample, as opposed to absolute concentrations. The concentrations of dicarboxylic acids measured in the Murchison meteorite by Martins et al. (2006) are lower than those measured by Shimoyama and Shigematsu (1994) and Peltzer et al. (1984), particularly for glutaric (C5) and adipic (C6) acids. For example, the ratios of succinic:glutaric acid are 4.0, 2.0, and 42.1 as measured by Shimoyama and Shigematsu (1994), Peltzer et al. (1984), and Martins et al. (2006), respectively, and the ratios of succinic:adipic acid reported by the same workers are 16.7, 10.4, and 29.3. An earlier version of this publication was completed prior to the publication of Martins et al. (2006). Their data have subsequently been included in this version. Although inclusion of the Martins et al. data increases the scatter in the plots presented in Figure 1, the overall results and conclusions remain unchanged.

If the data for the C4 through C9 acids obtained from the CM2 chondrites do reflect metastable equilibrium, the oxidation state of the equilibration (Figure 3) may be calculated based on the $\log a_{\text{C}_y\text{H}_{(2y-2)}\text{O}_4}$ -intercepts. Intercept values (Table 3) and equilibrium constants (Table 2) for the C4 through C6 acids were selected because these compounds were present in all four of the studies that examined CM2 chondrites. Although only three reactions are used to calculate the intercept values, each reaction is considered

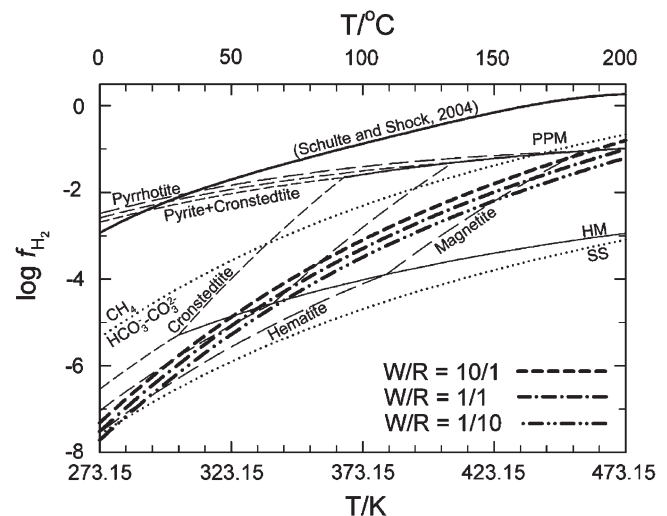


Figure 3. Calculation of $\log f_{\text{H}_2}$ in this work at water:rock ratios of 1:1 (dash-dot line), 10:1 (dotted line), and 1:10 (dash-dot-dot) based on the y -intercept of graphs indicating metastable equilibrium between succinic, glutaric, and adipic acids at 298.15–473.15 K. The Hematite-Magnetite (HM) and Pyrite-Pyrrhotite-Magnetite (PPM) buffers are shown extending from the right side of the diagram to the dashed line on the left side. The dashed lines define the cronstedtite stability field when $\log a_{\text{SiO}_2}$ is equal to quartz saturation minus 0.5 log units (short dashed line), quartz saturation (medium dashed line), and quartz saturation plus 0.5 log units (long dashed line). The short dashed lines designate the $\text{CH}_4(\text{aq})-\text{HCO}_3^-(\text{aq})-\text{CO}_3^{2-}(\text{aq})$ and $\text{H}_2\text{S}(\text{aq})-\text{SO}_4^{2-}(\text{aq})-\text{HS}^-(\text{aq})$ triple points. Also shown is the $\log f_{\text{H}_2}-T$ path generated during the coupled mineral alteration/organic synthesis simulation by Schulte and Shock (2004).

Table 3. Intercept values calculated by minimizing variance of log activity plots for slopes fixed by reaction stoichiometry

Water/rock	Reaction					
	A	B	C	D	E	F
1/1	-0.26	-0.66	0.32	-0.50	0.99	0.60
10/1	-0.46	-0.99	0.57	-0.67	1.49	0.80
1/10	-0.06	-0.33	0.07	-0.34	0.49	0.40

Reactions:

- (A) $5\text{C}_4\text{H}_6\text{O}_4 + 5\text{H}_2 = 4\text{C}_5\text{H}_8\text{O}_4 + 4\text{H}_2\text{O}$;
 (B) $6\text{C}_4\text{H}_6\text{O}_4 + 10\text{H}_2 = 4\text{C}_6\text{H}_{10}\text{O}_4 + 8\text{H}_2\text{O}$;
 (C) $4\text{C}_5\text{H}_8\text{O}_4 + 4\text{H}_2\text{O} = 5\text{C}_4\text{H}_6\text{O}_4 + 5\text{H}_2$;
 (D) $6\text{C}_5\text{H}_8\text{O}_4 + 5\text{H}_2 = 5\text{C}_6\text{H}_{10}\text{O}_4 + 4\text{H}_2\text{O}$;
 (E) $4\text{C}_6\text{H}_{10}\text{O}_4 + 8\text{H}_2\text{O} = 6\text{C}_4\text{H}_6\text{O}_4 + 10\text{H}_2$;
 (F) $5\text{C}_6\text{H}_{10}\text{O}_4 + 4\text{H}_2\text{O} = 6\text{C}_5\text{H}_8\text{O}_4 + 5\text{H}_2$.

twice. A first-order linear regression estimates the slope and intercept by minimizing the squared errors between the actual and predicted values of the dependent variable. In this exercise the slope is fixed by the reaction stoichiometry. Swapping the independent and dependent variables will invert the slope, producing a different intercept, a new value of s^2/N , and possibly a different $\log f_{\text{H}_2}$ value.

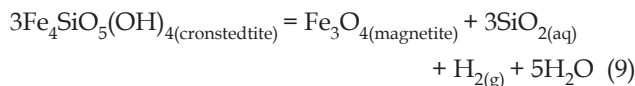
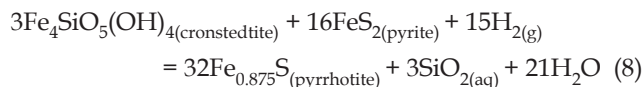
The calculated $\log f_{\text{H}_2}$ values agree within 0.3 log units at 298.15 K with the maximum discrepancy increasing to 0.6 log units at 473.15 K. The water:rock ratio selected for calculation of dicarboxylic acid activity will affect the $\log a_{\text{C}_2\text{H}_2(\text{Y}-2)\text{O}_4}$ -intercept value and thus the calculation of $\log f_{\text{H}_2}$. Varying water:rock ratios by a factor of 100 results in only a limited variation (0.4 log units) of calculated $\log f_{\text{H}_2}$ values (Figure 3).

4. Discussion

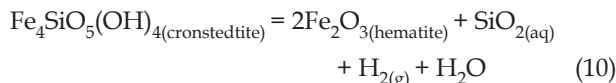
Do the $\log f_{\text{H}_2}$ values calculated here accurately reflect the oxidation state of fluids during any stage of aqueous alteration? Because the paragenesis of dicarboxylic acid synthesis and alteration is unknown, interpretation of the $\log f_{\text{H}_2}$ values calculated here is dependent on an interpretation of the thermal history of the carbonaceous chondrite parent body. Precise estimates of the temperature history during aqueous alteration are lacking. Measurements of the $\delta^{18}\text{O}$ values of bulk phyllosilicates and carbonates from the Murchison meteorite have been interpreted as evidence that oxygen isotope exchange occurred at low temperatures (273.15–298.15 K) (Clayton and Mayeda, 1999). These measurements are difficult ones to make: the individual mineral grains are small and calcite and dolomite CO_2 is recovered by sequential acidification of bulk samples (e.g. Benedix et al., 2003). Although the $\Delta^{17}\text{O}$ values of Murchison carbonates and the water extracted from Murchison phyllosilicates are all similar ($\sim -0.8\text{‰}$), the value reported for the Murchison matrix is much lower (-1.8‰) (Franchi et al., 2001). If the values for the Murchison matrix are representative of the phyllosilicate fraction then the phyllosilicate fraction did not form in isotopic equilibrium with either the carbonates or the phyllosilicate structural or interlayer water (Franchi et al., 2001). Baker et al. (2002) calculate an alteration temperature of ~ 353 K using the $\delta^{18}\text{O}$ values of the structural and interlayer water from phyllosil-

icates and calcite from Murchison ($\Delta^{18}\text{O}_{\text{calcite-water}} = 19\text{‰}$). Measurements of the Δ_{47} anomaly (Ghosh et al., 2006) of three splits of Murchison carbonate yield alteration temperatures ranging from 293 to 306 K, whereas temperatures up to 344 K are obtained from analyses of carbonate in the Cold Bokkeveld meteorite (Guo and Eiler, 2007). Temperatures on asteroids have been interpreted to be 378–398 K (Hayatsu and Anders, 1981), with other workers suggesting that alteration occurred at temperatures as high as 403 K (Bunch and Chang, 1980) or 423 K (Zolensky et al., 1993). Modeling exercises have utilized temperatures ranging from 273.15 to 373.15 K (Cohen and Coker, 2000), 273.15 to 473.15 K (Schulte and Shock, 2004), and 274.15–423.15 K (Zolensky et al., 1989). Aqueous alteration on the Tagish Lake parent body was estimated to occur at temperatures < 373.15 K (Zolensky et al., 2002).

We have attempted to relate the values of $\log f_{\text{H}_2}$ calculated in this study using organic species to the inorganic chemistry of the meteorites using Figure 3. Figure 3 was drawn for a temperature range of 273.15–473.15 K. The hematite-magnetite (HM) and pyrite-pyrrhotite-magnetite (PPM) buffers are designated by the solid lines that begin at the right side of the diagram and terminate against the cronstedtite field on the left of the diagram. Cronstedtite is one of the more abundant alteration minerals in CM2 chondrites. It is stable at reducing conditions, low temperatures and activities of aqueous silica that approach quartz saturation (e.g. Figure 2 in Dyl et al., 2006). The three reactions that constrain cronstedtite stability on Figure 3 are



and



Reactions (8), (9), and (10) have been plotted in $\log f_{\text{H}_2}$ - T space (Figure 3) by fixing the activity of aqueous silica to 0.5 log units below quartz saturation (the short dashed line), quartz saturation, and 0.5 log units above (the long dashed line) quartz saturation. Although silica supersaturation will not occur under equilibrium conditions, amorphous silica is produced during computational modeling of the aqueous alteration of CM chondrites (Dyl et al., 2006). Moreover, short-term, high-temperature (up to 473.15 K) aqueous alteration experiments using the Allende meteorite have produced amorphous, silica-rich phases (Jones and Brearley, 2006). Cronstedtite is typically stable only at low temperatures. Nucleation of cronstedtite at temperatures above 333 K requires kinetic inhibition of minnesotaite, a phase not typically reported in CM chondrites. The reaction path calculation of Schulte and Shock (2004) does not produce cronstedtite above ~ 323 K but cronstedtite-bearing alteration mineral assemblages are preserved at temperatures up to 373 K.

Table 4. $\log f_{\text{H}_2}$ values calculated based on water/rock = 1/1

T (K)	$\log f_{\text{H}_2}$ values calculated from reactions A-F					
	A	B	C	D	E	F
298.15	-6.3	-6.2	-6.3	-6.0	-6.2	-6.0
348.15	-4.3	-4.1	-4.3	-3.9	-4.1	-3.9
373.15	-3.5	-3.3	-3.5	-3.1	-3.3	-3.1
423.15	-2.2	-2.0	-2.2	-1.7	-2.0	-1.7
473.15	-1.3	-1.0	-1.3	-0.7	-1.0	-0.7

Reactions:

- (A) $5\text{C}_4\text{H}_6\text{O}_4 + 5\text{H}_2 = 4\text{C}_5\text{H}_8\text{O}_4 + 4\text{H}_2\text{O}$;
 (B) $6\text{C}_4\text{H}_6\text{O}_4 + 10\text{H}_2 = 4\text{C}_6\text{H}_{10}\text{O}_4 + 8\text{H}_2\text{O}$;
 (C) $4\text{C}_5\text{H}_8\text{O}_4 + 4\text{H}_2\text{O} = 5\text{C}_4\text{H}_6\text{O}_4 + 5\text{H}_2$;
 (D) $6\text{C}_5\text{H}_8\text{O}_4 + 5\text{H}_2 = 5\text{C}_6\text{H}_{10}\text{O}_4 + 4\text{H}_2\text{O}$;
 (E) $4\text{C}_6\text{H}_{10}\text{O}_4 + 8\text{H}_2\text{O} = 6\text{C}_4\text{H}_6\text{O}_4 + 10\text{H}_2$;
 (F) $5\text{C}_6\text{H}_{10}\text{O}_4 + 4\text{H}_2\text{O} = 6\text{C}_5\text{H}_8\text{O}_4 + 5\text{H}_2$.

The $\log f_{\text{H}_2}$ values calculated in this study are consistent with those estimated based on previous mineralogical examinations of carbonaceous chondrites (e.g. Zolensky et al., 1993) but cover a narrower range of $\log f_{\text{H}_2}$ - T space. The curves calculated from this study (Figure 3) are the average values of $\log f_{\text{H}_2}$ as a function of temperature for the six reactions examined in Table 4 calculated for three different values of water:rock ratio. The Schulte and Shock (2004) path in Figure 3 is the $\log f_{\text{H}_2}$ - T path for their coupled mineral alteration/organic synthesis simulation from 273.15 to 473.15 K. If the path developed in the Schulte and Shock (2004) simulation is an accurate approximation of the prograde alteration on the parent body, then the values of $\log f_{\text{H}_2}$ calculated in this study may record a later, retrograde, more oxidizing, period of aqueous alteration. The values of $\log f_{\text{H}_2}$ calculated here are, however, generally consistent with the mineralogy of the Murchison and Y-791118 meteorites. The dicarboxylic acid species could therefore have reached metastable equilibrium at temperatures approaching 353 K in contact with cronstedtite + magnetite and in an aqueous solution in which the activity of aqueous silica approached quartz saturation (Figure 3). The maximum temperature range of equilibrium with cronstedtite ranges from 300 to 355 K if a water:rock ratio of 10:1 and quartz saturation are assumed.

The values calculated in this study are more oxidizing than those presented in a recent examination of the $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ values of carbonates from the Murchison, Murray and Cold Bokkevoeld meteorites. Guo and Eiler (2007) interpreted a negative correlation between the carbonate $\delta^{13}\text{C}$ values and the calculated $\delta^{18}\text{O}$ values for the alteration water as evidence for reduction of aqueous carbonate to methane during alteration with subsequent loss of the methane. Guo and Eiler (2007) obtained an Eh value of -0.67 V ($\log f_{\text{H}_2} = -2.6$) and a pH value of 12.5 by modeling a fluid in equilibrium with a CM chondrite alteration assemblage at 301 K. Methane loss can occur under conditions slightly more oxidizing than those defined by the trace of the $\text{CH}_4(\text{aq}) - \text{HCO}_3^-(\text{aq}) - \text{CO}_3^{2-}(\text{aq})$ triple point (Figure 3) depending on total dissolved carbon, fluid salinity, and pressure. Again, if the values obtained in this study

are accurate, then they must reflect a later, more oxidizing period of alteration. Methane loss could have been a significant mechanism for fluid oxidation if the loss of methane was rapid relative to heterogeneous equilibrium.

Progressively more oxidizing conditions during aqueous alteration are indicated by the presence of calcite and sulfate veins proximal to more reduced mineral assemblages (Kerridge and Bunch, 1979). Although the $\Delta^{17}\text{O}$ values of sulfate from CM2 chondrites deviate from the terrestrial fractionation line, the values obtained from Murchison sulfate do not match the $\Delta^{17}\text{O}$ values of the phyllosilicate matrix, phyllosilicate structural or interlayer water, or carbonates (cf. [Baker et al., 2002; Airieau et al., 2005]). Nor are the distribution of dicarboxylic acid species consistent with metastable equilibrium during a sulfide oxidizing event. The $\log f_{\text{H}_2}$ values of the $\text{H}_2\text{S}(\text{aq}) - \text{SO}_4^{2-}(\text{aq}) - \text{HS}^-(\text{aq})$ triple point are plotted in Figure 3 for reference. The dicarboxylic acid species could only reach metastable equilibrium with a sulfate fluid at temperatures approaching 273.15 K.

It must be noted that specific ratios in the abundances of organic species, interpreted as evidence of metastable equilibrium (Shock, 1988; Shock, 1989; Shock, 1990), may originate from other processes. One example is the generation of monocarboxylic acids during the thermal maturation of kerogen (Bell and Palmer, 1994). Extending this argument to the carbonaceous chondrite samples, the soluble cyclic compounds in Murchison have been suggested to originate from macromolecular material (Sephton et al., 1998). Pyrolysis of the Yamato-791198 carbonaceous chondrite yielded an array of heterocyclic compounds, saturated alkanes of 9-14 carbon atoms, propanenitrile, and dimethylsulfide from the macromolecular material (Komiya et al., 1993). Carboxylic acids, however, were not identified and evidence for the synthesis of carboxylic acids during pyrolysis of chondrite macromolecular organic material is lacking.

5. Conclusions

The concentrations of C4-C9 saturated dicarboxylic acids in carbonaceous chondrite samples are correlated in a way that suggests that the acids were in metastable equilibrium during aqueous alteration on the meteorite parent body. Because the slopes of these relations are fixed by the stoichiometry of the reactions relating the species of interest, the oxidation state can be retrieved. The linearity of the plots based on samples of different carbonaceous chondrites suggests similar thermal and redox histories among the CM2 carbonaceous chondrites sampled and that the Murchison and Y-791118 meteorites were derived from the same parent body. Calculation of $\log f_{\text{H}_2}$ values as a function of temperatures is possible by minimizing the variance on the plots on which the slope is fixed by the stoichiometry of the reaction indicating metastable equilibrium. Based on metastable equilibrium among succinic (C4), glutaric (C5), and adipic (C6) acids and a water:rock ratio of 1:1 during aqueous alteration, values of $\log f_{\text{H}_2}$ calculated range from -6.2 at 298 K to -3.3 at 373 K. The values of $\log f_{\text{H}_2}$ calculated are consistent with those proposed for hydrothermal alteration of carbonaceous chon-

drites based on petrographic studies but more oxidizing than values proposed for methane-producing systems. The results obtained here support the application of metastable equilibrium to problems of organic synthesis in aqueous systems.

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